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## Effects of acetylene at low concentrations on nitrification, mineralization and microbial biomass nitrogen concentrations in forest soils

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Temperate forest surface soils at the varying distances from main trunks (e.g., Pinus koraiensis and Quercus mongolica) were used to study the effects of acetylene (C<sub>2</sub>H<sub>2</sub>) at low concentrations on nitrification, mineralization and microbial biomass N concentrations of the soils, and to assess the contribution of heterotrophic nitrification to nitrous oxide (N2O) emissions from soils. The use of acetylene at partial pressures within a range from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas gave a significant decrease in N<sub>2</sub>O emission at soil moisture of c. 45% water-filled porosity space, and the decrease was almost the same in each soil after exposure of C<sub>2</sub>H<sub>2</sub> at low concentrations. Heterotrophic nitrification could account for 21%-48% of total N<sub>2</sub>O emission from each soil; the contribution would increase with increasing distances from the Pinus koraiensis trunks rather than from the Quercus mongolica trunks. Under the experimental conditions, the use of C<sub>2</sub>H<sub>2</sub> at low concentrations showed no significant influence on soil microbial biomass N, net N mineralization and microbial respiration. However, 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas could reduce carbon dioxide (CO2) emissions from soils. According to the rapid consumption of 10 Pa C<sub>2</sub>H<sub>2</sub> by forest soils and convenience for laboratory incubations, 50 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas can be used to study the origin of N2O emissions from forest soils under aerobic conditions and the key associated driving mechanisms. The N<sub>2</sub>O and CO<sub>2</sub> emissions from the soils at the same distances from the Quercus mongolica trunks were larger than those from the Pinus koraiensis trunks, and both emissions decreased as the distances from trunks increased. The stepwise regression analysis showed that 95% of the variability in soil CO2 emissions could be accounted for by the concentrations of soil total C and water soluble organic C and soil pH, and that 72% of the variability in soil N₂O emissions could be accounted for by the concentrations of soil total N, exchangeable NH₄⁺-N and microbial biomass N and 25% of the variability in heterotrophic nitrification by the soil microbial biomass N concentration. The emissions of N2O and CO2 from forest soils after exposure of C2H2 at low concentrations were positively related to the net nitrification of the soils.

heterotrophic nitrification, forest soil, acetylene, mineralization, microbial biomass nitrogen

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Low doses of acetylene (C<sub>2</sub>H<sub>2</sub>) (normally within a range of 10 to 100 Pa) inhibited autotrophic NH<sub>4</sub> oxidation, and then NH<sub>4</sub> concentration in soil was increased after exposure of  $C_2H_2^{[1-3]}$ . Due to microbial preferential utilization for NH<sub>4</sub>, this phenomenon may result in an increase in microbial N immobilization in soil<sup>[4]</sup>, which was accompanied by the decrease in soil mineral N concentration. As a result, the change in concentrations of soil N forms due to addition of C<sub>2</sub>H<sub>2</sub> is likely to affect the responses of nitrous oxide (N2O) emissions from soils to addition of C<sub>2</sub>H<sub>2</sub> at low concentrations. Hatch et al. [5] showed that the use of acetylene at larger than 10 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas tended to increase soil microbial N immobilization. In addition, C2H2 can act as a carbon source to heterotrophic soil microorganisms [6], thereby increasing soil microbial N concentration at partial pressures larger than 100 Pa C<sub>2</sub>H<sub>2</sub><sup>[7]</sup>. To date, there is limited knowledge about the effects of C<sub>2</sub>H<sub>2</sub> exposure on soil microbial biomass N concentration. This lack would not be beneficial to further studying soil N transformations and the key associated driving mechanisms by using inhibition of  $C_2H_2$ .

Earlier few studies have shown that there are reduced rates of mineralization in grassland soils at concentrations of added C<sub>2</sub>H<sub>2</sub> from 500 to 10000 Pa C<sub>2</sub>H<sub>2</sub>, and that 10% C<sub>2</sub>H<sub>2</sub> in headspace gas can result in an apparent net immobilization of mineral N<sup>[8]</sup>. However, we cannot confirm whether the exposure of C<sub>2</sub>H<sub>2</sub> at low concentrations affects the change in soil mineral N production. Due to the limitations of measurement, at present there is limited knowledge with regard to heterotrophic nitrification and its relationships to N transformations and N<sub>2</sub>O production in acid forest soils<sup>[9]</sup>. Differences in stem flow of different trees normally result in the change in properties of the soils around trunks, which is likely to affect the origin of N<sub>2</sub>O emission from soil and its spatial properties under forest stands. Hence, it is useful and timely to understand the effects of C<sub>2</sub>H<sub>2</sub> at low concentrations on microbial N immobilization and net mineralization in soil.

The objectives of this present work were to study the effects of  $C_2H_2$  at low concentrations on nitrification, mineralization and microbial N immobilization in temperate forest surface soils at the varying distances (0, 50, and 100 cm) from coniferous and hardwood trunks. Of addition interest are the contributions of heterotrophic nitrification and the varying distances from trunks to  $N_2O$  production in soil. The results would promote one

to study the origin of N<sub>2</sub>O emissions from acid soils under upland forest ecosystems and the key associated driving mechanisms.

#### Materials and methods

## 1.1 Soil sampling and measurement of soil properties

Temperate forest surface soils under the Korean pine and broadleaf mixed forest (Pinus koraiensis mainly mixed with hardwood trees such as Quercus mongolica, >200 years, altitude 740 m) were sampled near the Research Station of Changbai Mountain Forestry Ecology (128°28′E, 42°24′N). Annual mean temperature is c.  $2.8^{\circ}$ C and precipitation c. 750 mm at the bottom of the mountain. Five individual coniferous and broadleaf trees (e.g., Pinus koraiensis and Ouercus mongolica) with a broad diameter of c. 80 cm each and with no trees within the range of 10 to 15 m from trunks were respectively selected in July 2006, and at the varying distances (0, 50, and 100 cm) from trunks, forest surface soils to a depth of 7 cm, after removal of litter and humus layers, were sampled by hammering the cores into the soil, and bulked as one sample. Hence, in total 30 bulk soil samples were taken. All samples were kept separately in airtight plastic bags and rapidly transported to the laboratory. The fresh moist soils were sieved (2 mm) to remove small stones and roots, and stored in the dark at 4°C for use.

The forest soil belongs to Andisols (Food and Agriculture Organization Soil Classification), and the depth of the A-horizon in soil profile is approximately 10 cm. The properties of the soils at the various depths beneath the forest stand were reported by Xu et al. [10]. Duplicate soils were dried at 105°C for 24 h to determine moisture content. Total C and N concentrations in soil samples were measured with a CN analyzer (MT-700 with an Auto Sampler MTA-600, Yanaco, Kyoto, Japan). Fresh soil pH (soil/water, 1/2.5, w/w) was measured with a portable pH meter. Exchangeable NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N concentrations were extracted by shaking 5.0-g fresh soil with 25 mL of 1 mol/L KCl solution for 60 min on an end-over-end shaker. The slurry was then filtered into 50-mL plastic bottles, and NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N concentrations in the extracts were measured colorimetrically via the nitroprusside and hydrazine-reduction methods, respectively [11]. Forest soils (5.0 g) were extracted by shaking with 25 mL of deionized water for 30 min on an

end-over-end shaker. The suspensions were centrifuged at 6400 g for 5 min and then filtered into 50-mL plastic bottles via cellulose-acetate membrane filters (0.45 µm pore size). Soluble organic C and total N concentrations in the soil extracts were measured using a TOC/TNanalyzer (Shimadzu TOC-V<sub>CSH</sub>/TN, Kyoto, Japan). Concentrations of soluble soil organic N (DON) were considered the differences between total N and mineral N (NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N) concentrations in the soil extracts. Soil microbial C and N concentrations were measured using the chloroform fumigation-extraction method<sup>[12]</sup>. Fresh soils (5.0 g) with and without chloroform fumigation were extracted with 25 mL of 0.5 mol/L K<sub>2</sub>SO<sub>4</sub> solution, as above, and soluble organic C and total N in the extracts were measured using a TOC/TNanalyzer (Shimadzu TOC-V<sub>CSH</sub>/TN, Kyoto, Japan). The main soil properties at the varying distances from main trunks are presented in Table 1.

#### 1.2 Incubation experiment

All forest soil samples taken from the varying distances from trunks were used to study the change in soil microbial biomass N and mineral N concentrations and in  $N_2O$  and  $CO_2$  emissions from the soils after exposure of  $C_2H_2$  at low concentrations. For each of soil samples, water was added to naturally settled soils (20.0 g) in 250-mL glass bottle to achieve a range of WFPS from 0.40—0.45 m³ water m⁻³ pore space. The level of the soil in each bottle was marked prior to incubation and the volume of the soil was measured via water following the removal of soil. The WFPS was calculated according to the equation reported by Franzluebbers [13]. At the start of the incubation, a series of acetone-free  $C_2H_2$  (approximately 3%  $C_2H_2$  in  $N_2$ ,  $\nu/\nu$ ) volume was injected into each bottle through a stopcock connected to the needle

inserted through the stoppers using air-tight syringes to give concentrations of 10, 50 and 100 Pa C<sub>2</sub>H<sub>2</sub>; the control with C<sub>2</sub>H<sub>2</sub> addition was incubated. The batch incubation was done at 25°C in the dark for 10 days under oxic conditions. Twenty-five milliliter of headspace gases was sampled from each bottle using 50-mL air-tight syringes 3, 5, 7 and 10 days after C<sub>2</sub>H<sub>2</sub> addition and then analyzed for concentrations of CO2, N2O and C<sub>2</sub>H<sub>2</sub> using a gas chromatograph. After sampling, the rubber stoppers were removed from the bottles for 2 h to refresh the air, and then they were closed again. A series of acetone-free C<sub>2</sub>H<sub>2</sub> volume was again injected into each bottle, as above, to receive concentrations of 10, 50 and 100 Pa C<sub>2</sub>H<sub>2</sub>. Preliminary studies have indicated that 10 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas is completely consumed by soil microorganisms within 24 h. Thus, the presence of 10 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas can remain through daily injection of standard C<sub>2</sub>H<sub>2</sub> into each bottle. At the end of the 10-day oxic incubation, 5-g and 10-g subsamples from each treatment were respectively used to measure the concentrations of soil mineral N (NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N) and microbial biomass C and N, as above.

### 1.3 Measurement of CO<sub>2</sub>, N<sub>2</sub>O and C<sub>2</sub>H<sub>2</sub> concentrations

According to the methods reported by Wang et al. [14], concentrations of CO<sub>2</sub>, N<sub>2</sub>O and C<sub>2</sub>H<sub>2</sub> in headspace gases were respectively quantified with a modified gas chromatograph (Agilent 5890) equipped with a flame ionization detector (FID) and with an electron capture detector (ECD). Carbon dioxide, N<sub>2</sub>O and C<sub>2</sub>H<sub>2</sub> were severally separated by one stainless steel column (2 m length and 2.2 mm i.d.) that was packed with 50–80 mesh Porapack Q, afterwards hydrogen reduced CO<sub>2</sub> to CH<sub>4</sub> in a Nickel catalytic converter at 375°C, and the

Table 1 Main properties of the soils at the varying distances from trunk
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Distances from	TC	TN	– C/N	DOC	DON	Microbial C	Microbial N	NH <sub>4</sub> -N	NO <sub>3</sub> -N	рН
trunks (cm)	(mg g <sup>-1</sup> dry soil)		= C/N	(μg g <sup>-1</sup> dry soil)						
Broadleaf fore	st soil									
0	129.7(7.5)	11.5(1.6)	11.7(0.8)	621(118)	119(10)	2338(171)	573(27)	28.0(4.6)	2.7(0.8)	6.3(0.1)
50	78.8(3.6)	6.8(0.4)	11.7(0.3)	299(44)	69(6)	1939(187)	441(42)	16.3(3.1)	1.1(0.1)	5.6(0.1)
100	72.8(10.1)	6.3(0.8)	11.4(0.3)	252(30)	59(6)	1920(266)	429(62)	11.4(2.9)	0.9(0.1)	5.4(0.1)
Coniferous for	est soil									
0	71.6(17.8)	5.4(0.9)	12.8(0.8)	512(201)	71(15)	1620(261)	369(73)	17.3(5.3)	2.4(0.4)	5.4(0.1)
50	50.6(6.2)	4.1(0.5)	12.4(0.5)	245(38)	46(4)	1239(167)	281(36)	7.9(1.5)	0.9(0.1)	5.5(0.1)
100	46.7(5.1)	3.9(0.4)	12.1(0.4)	192(11)	45(4)	1108(194)	245(45)	7.1(1.6)	0.8(0.1)	5.5(0.1)
LSD <sub>0.05</sub>	28.1	2.7	1.6	289	25	619	145	10.1	1.1	0.2

a) Values are the means of five replicates, and standard errors are shown in parenthesis. Least significant differences (LSD; 5%) were used to assess differences in the soil properties at the varying distances from trunks. Broadleaf tree, *Pinus koraiensis*; Coniferous tree, *Quercus mongolica*. The same below.

CH<sub>4</sub> and C<sub>2</sub>H<sub>2</sub> were severally detected by FID. The oven was operated at 55 °C, and the FID at 200 °C (ECD at 330 °C for N<sub>2</sub>O), respectively, with N<sub>2</sub> as carrier gas at a flow rate of 20 cm<sup>3</sup> min<sup>-1</sup>. The detector responses were calibrated using certified gas standards, which contain 4.0 mL L<sup>-1</sup> CO<sub>2</sub> in N<sub>2</sub>, 0.37  $\mu$ L L<sup>-1</sup> N<sub>2</sub>O in N<sub>2</sub>, and 100  $\mu$ L L<sup>-1</sup> C<sub>2</sub>H<sub>2</sub> in N<sub>2</sub>, respectively.

#### 1.4 Calculation and statistical analysis

All data were transformed on an oven-dried weight basis. Rates of N<sub>2</sub>O and CO<sub>2</sub> emissions from soils were calculated from the increase in gas concentration with time, and were expressed in pmol N<sub>2</sub>O-N g<sup>-1</sup> dry soil d<sup>-1</sup> and umol CO<sub>2</sub>-C g<sup>-1</sup> dry soil d<sup>-1</sup>, respectively. The CO<sub>2</sub> concentrations in headspace gases excluded CO2 production due to the decomposition of acetylene by soil. Soil net N mineralization rates were calculated by the difference in soil NH<sub>4</sub><sup>+</sup>-N and NO<sub>3</sub><sup>-</sup>-N concentrations after and prior to the incubation divided by time, and were expressed in μg N g<sup>-1</sup> dry soil d<sup>-1</sup>. Concentrations of soil microbial biomass C and N were calculated by using the difference in salt-extractable organic C or N concentrations between fumigated and non-fumigated soil extracts divided by 0.45[15,16]. Means and standard errors of five replicates in each treatment were calculated. Least significant differences (LSD) were calculated at the 5% level to assess the differences in N<sub>2</sub>O and CO<sub>2</sub> emissions and in soil properties between treatments and soils. Stepwise regression analysis was performed by using SPSS software for Windows (version 13.0) to assess soil properties that can affect the emissions of  $N_2O$  and  $CO_2$  from soils and heterotrophic nitrification in forest soils.

#### 2 Results and discussion

## 2.1 $N_2O$ emissions from soils after exposure of $C_2H_2$ at low concentrations and heterotrophic nitrification

Table 2 shows the effects of addition of C<sub>2</sub>H<sub>2</sub> at low concentrations on N<sub>2</sub>O emissions from the soils at the varying distances from trunks, and the contribution of heterotrophic nitrification to the soil N<sub>2</sub>O emission. The N<sub>2</sub>O emission from the soil at the same distance from the hardwood trunks was larger than that from the coniferous trunks, and it decreased as the distances from trunks increased. The difference in soil N<sub>2</sub>O emissions was mainly ascribed to the status of nutrients and microbial biomass C and N concentrations in soil (Table 1). The results were in agreement with the findings reported by Ullah et al. (17), who showed that deciduous forest soils exhibited greater potential of N<sub>2</sub>O fluxes than did white pine forest soils in Eastern Canada.

In comparison with the control, the use of acetylene at partial pressures from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gases gave a significant decrease in N<sub>2</sub>O emission at soil moisture of c. 45% water-filled porosity space, and the decrease was the same in each soil after treatment with acetylene at low concentrations (Table 2). Heterotrophic nitrification could account for 21%—48% of total N<sub>2</sub>O emission from each soil; the contribution would increase with increasing distances from the coniferous trunks

**Table 2**  $N_2O$  emissions from the soils after exposure of  $C_2H_2$  at low concentrations and proportion of heterotrophic nitrification in the soils at the varying distances from trunks

Distance for a	N <sub>2</sub> O								
Distances from trunks (cm)	0 Pa C <sub>2</sub> H <sub>2</sub>	10 I	Pa C <sub>2</sub> H <sub>2</sub>	50 F	Pa C <sub>2</sub> H <sub>2</sub>	100	Pa C <sub>2</sub> H <sub>2</sub>	_ LSD <sub>0.05</sub> <sup>a)</sup>	LSD <sub>0.05</sub> b)
	N <sub>2</sub> O emission	N₂O emission	Proportion of heterotrophic nitrification	N <sub>2</sub> O emission	Proportion of heterotrophic nitrification	N₂O emission	Proportion of heterotrophic nitrification	_ L3D <sub>0.05</sub>	LOD <sub>0.05</sub>
Broadleaf fores	t soil								
0	133.5(23.4)	31.7(4.0)	0.26(0.04)	25.8(1.7)	0.21(0.02)	31.2(3.1)	0.27(0.06)	36.0	0.14
50	103.2(16.9)	29.1(4.7)	0.30(0.04)	26.8(6.4)	0.26(0.05)	21.0(3.5)	0.21(0.03)	28.4	0.12
100	104.8(18.5)	25.1(3.4)	0.30(0.10)	22.1(2.7)	0.30(0.13)	22.9(2.4)	0.30(0.12)	28.7	0.36
Coniferous fore	st soil								
0	109.9(5.6)	29.9(6.8)	0.27(0.05)	26.5(5.7)	0.24(0.05)	28.4(3.7)	0.26(0.05)	16.7	0.15
50	73.5(12.1)	22.5(4.1)	0.32(0.05)	15.9(1.2)	0.25(0.05)	18.8(3.8)	0.27(0.05)	20.0	0.15
100	66.6(12.4)	33.6(10.0)	0.48(0.08)	25.1(6.9)	0.37(0.06)	22.3(5.6)	0.34(0.04)	27.1	0.19
LSD <sub>0.05</sub> c)	46.2	17.4	0.19	13.7	0.20	10.6	0.19		

a) LSD (5%) for comparing differences in  $N_2O$  emissions from each soil after exposure of  $C_2H_2$  at low concentrations; b) LSD (5%) for comparing differences in proportions of heterotrophic nitrification to  $N_2O$  emission from each soil between treatments of  $C_2H_2$  at low concentrations; c) LSD (5%) for comparing differences in  $N_2O$  emissions and proportions of heterotrophic nitrification between soils after exposure of  $C_2H_2$  at the same concentration.

rather than from the hardwood trunks (Table 2). Xu and Inubushi<sup>[7]</sup> reported that heterotrophic nitrification among acid forest soils in central Japan accounted for the range from 37% to 76% of the total N<sub>2</sub>O production under the controlled aerobic conditions and was variable with forest stands, topographical slope and seasons. Hart et al. [18] also reported that heterotrophic nitrification accounted for 65%-72% of the gross nitrification in two coniferous forest stands of contrasting productivity and with different altitude, and that soil pH (3.9-5.4) and status of substances were important in controlling the significance of heterotrophic nitrification in forest soil. However, using <sup>15</sup>N pool dilution-enrichment approach in the laboratory, Barraclough and Puri observed that heterotrophic nitrification contributed ≤8% of the total gross nitrification rate in an acid woodland soil. There was usually low activity of autotrophic nitrifiers and the predominance of heterotrophic nitrification in acid forest soils at pH < 4.5, although there is information about autotrophic nitrifying bacteria that are capable of ammonium oxidation at pH 4<sup>[20]</sup>. Pedersan et al. [21] showed that the heterotrophic nitrification accounted for a larger contribution to total gross nitrification rate in coniferous forest soils after clear-cutting compared to that in undisturbed forest soils. Hence, the contribution of heterotrophic nitrification to soil total gross nitrification was dependent on kinds of forest vegetations, soil properties, and seasons.

Acetylene at low concentrations can be considered a particularly useful selective biochemical block of the autotrophic pathway<sup>[1,2]</sup>, which does not inhibit the het-

erotrophic pathway<sup>[3,22]</sup>. In order to effectively study heterotrophic nitrification and its contribution to  $N_2O$  production from acid forest soils, it is timely and important to understand the effects of acetylene at low concentrations on microbial N immobilization and net mineralization in soil.

# 2.2 Change in soil net mineralization and microbial biomass N concentrations after exposure of $C_2H_2$ at low concentrations

The net N mineralization rate in the soil at the same distance from the hardwood trunks was larger than that from the coniferous trunks, and it was larger at 0 cm than at 50 cm and 100 cm distances from each trunk (Table 3). The differences in soil net N mineralization rates can be also ascribed to the status of nutrients and microbial biomass N and C concentrations in soil (Table 1). In comparison with the control, the use of acetylene at partial pressures from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gases showed no significant influence on the net N mineralization in the soils at the varying distances from trunks. This was in agreement with the results reported by Ross et al. [23], who showed that the use of 100 Pa C<sub>2</sub>H<sub>2</sub> had no significant influence on net N mineralization in forest soils in Northeastern USA. However, Boer et al. [24] reported that the addition of acetylene at partial pressures from 10 to 1000 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas decreased net mineral N production in mineral layer rather than in organic layer of an acid oak-beech forest soil. Similarly, Hatch et al. [8] reported that the use of C<sub>2</sub>H<sub>2</sub> at elevated concentrations from 1000 to 10000 Pa C<sub>2</sub>H<sub>2</sub> significantly decreased net mineral N production

Table 3 Net N mineralization and microbial biomass N concentrations after exposure of C<sub>2</sub>H<sub>2</sub> at low concentrations in the soils at the varying distances from trunks

	Soil r	net N mineraliz	ation (µg N g <sup>-1</sup> dr	y soil d <sup>-1</sup> ) and	microbial biomas	s N concentrati	ons (µg N g <sup>-1</sup> dry	soil)				
Distances	after exposure of $C_2H_2$ at low concentrations											
from trunks	0 Pa C₂H₂		10 Pa C <sub>2</sub> H <sub>2</sub>		50 Pa C	C <sub>2</sub> H <sub>2</sub>	100 Pa	LSD <sub>0.05</sub> <sup>a)</sup>	LSD <sub>0.05</sub> b)			
(cm)	Net N	Microbial	Net N	Microbial	Net N	Microbial	Net N	Microbial	=			
	mineralization	N	mineralization	N	mineralization	N	mineralization	N				
Broadleaf	forest soil											
0	7.83(0.90)	573.1(26.8)	7.92(1.14)	597.2(29.4)	7.75(0.94)	594.1(37.1)	7.65(0.79)	613.3(39.0)	2.85	100.4		
50	2.77(0.22)	441.2(41.7)	2.69(0.22)	469.5(47.8)	2.58(0.32)	458.0(42.2)	2.93(0.29)	465.3(43.2)	0.81	131.3		
100	3.51(0.47)	429.0(61.7)	3.32(0.42)	446.2(63.6)	3.50(0.24)	456.8(69.9)	3.32(0.33)	448.5(66.5)	1.3	196.4		
Coniferous	s forest soil											
0	2.63(0.66)	369.3(72.5)	2.44(0.61)	363.0(67.4)	2.69(0.66)	360.3(67.5)	2.72(0.69)	367.5(67.0)	1.54	205.8		
50	2.10(0.17)	281.9(36.2)	1.90(0.34)	284.5(40.6)	1.80(0.26)	279.4(46.7)	1.80(0.27)	284.7(39.5)	0.79	122.7		
100	1.68(0.48)	245.5(44.7)	1.69(0.30)	253.6(40.8)	1.65(0.34)	251.6(42.9)	1.74(0.30)	251.4(43.4)	1.08	128.8		
LSD <sub>0.05</sub> c)	1.51	145.2	1.67	146.1	1.46	153.7	1.34	149.6				

a) LSD (5%) for comparing differences in net N mineralization in each soil after exposure of  $C_2H_2$  at low concentrations; b) LSD (5%) for comparing differences in microbial biomass N concentrations in each soil after exposure of  $C_2H_2$  at low concentrations; c) LSD (5%) for comparing differences in net N mineralization and microbial biomass N concentrations between soils after exposure of  $C_2H_2$  at the same concentration.

in soil and the inhibition was associated with initial soil  $NO_3^-$  concentration. Hence, the effects of  $C_2H_2$  addition on soil net N mineralization were dependent on the concentrations and decomposition of  $C_2H_2$  and soil properties such as  $NO_3^-$ N concentration.

Soil microbial biomass N concentration at the same distance from the hardwood trunks was much larger than that from the coniferous trunks (P < 0.05), and it decreased as the distances from trunks increased. The difference in soil microbial biomass N concentrations could be accounted for by the relatively large concentrations of organic C and water soluble organic C in the soils under hardwood trees rather than under coniferous trees (Table 1). In comparison with the control, the use of acetylene at partial pressures from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gases showed no significant influence on microbial biomass N concentration in soil (Table 3). However, Hatch et al. [5] showed that the addition of acetylene at partial pressures larger than 10 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas tended to increase microbial N immobilization in fertilized and unfertilized grass soils, but did not influence microbial N concentration in unfertilized grass and clover soil. The oxidizing capacity of the nitrifiers was recognized as one of important factors controlling the increase in soil microbial N immobilization<sup>[5]</sup>. Xu and Inubushi [7] reported that the addition of C<sub>2</sub>H<sub>2</sub> at partial pressures within a range from 500 to 10000 Pa C<sub>2</sub>H<sub>2</sub> in headspace gases significantly increased microbial N immobilization in soil. Furthermore, the stimulating effect of C<sub>2</sub>H<sub>2</sub> on the increase in soil microbial N immobilization would remain a saturation status after exposure of leveled C<sub>2</sub>H<sub>2</sub> concentrations<sup>[7]</sup>. Hence, the effects of C<sub>2</sub>H<sub>2</sub> on soil microbial N immobilization were dependent on the concentrations of C<sub>2</sub>H<sub>2</sub>, kinds of soils, the oxidizing capacity of the nitrifiers, and the other experimental conditions.

## 2.3 Change in soil respiration after exposure of $C_2H_2$ at low concentrations

Soil respiration has been considered one of main methods describing apparent microbial activity. According to the measurement of  $CO_2$  emissions from soils, the present experiment can be used to assess the effects of  $C_2H_2$  at low concentrations on microbial respiration in forest soils. The results and combination with the responses of soil microbial N immobilization and net mineralization to  $C_2H_2$  exposure would promote one to study the origin of  $N_2O$  emissions from acid forest soils and the key as-

sociated driving mechanisms.

The addition of acetylene at partial pressures from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gases showed no significant influence on microbial respiration in forest soils, but the use of 100 Pa C<sub>2</sub>H<sub>2</sub> tended to decrease soil respiration (Table 4). The CO<sub>2</sub> emission from the soil at the same distance from the hardwood trunks was larger than that from the coniferous trunks, and it decreased as the distances from trunks increased (Table 4). This trend paralleled with the change in N<sub>2</sub>O emissions, net mineralization and microbial biomass N concentrations of the soils with and without addition of acetylene (Tables 2 and 3). Herrmann et al. [25] showed that brief exposure to low atmospheric C<sub>2</sub>H<sub>2</sub> concentrations in headspace gas (0.1-1.0 kPa) did not significantly affect C mineralization rates when nitrification was completely inhibited. However, in a wheat soil from Hebei Province, Li et al. [26] reported that the use of 2000 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas reduced the respiration of soil heterotrophic microorganisms by 50%. Hence, the use of acetylene at low concentrations does not affect the activity of soil microbial respiration, but the activity can be inhibited by leveled C<sub>2</sub>H<sub>2</sub> concentrations. The functioning of the use of acetylene in inhibiting soil carbon mineralization is variable with soil properties and experimental conditions.

# 2.4 Soil properties affecting $N_2O$ and $CO_2$ emissions from the soils at the varying distances from trunks

Table 5 shows the coefficients of correlation between  $N_2O$  and  $CO_2$  emissions from the soils in the absence of acetylene and soil properties. The  $N_2O$  and  $CO_2$  emissions from the soils without  $C_2H_2$  addition were significantly related to the concentrations of soil total C and N, water soluble organic N, and microbial biomass C and N, respectively. Heterotrophic  $N_2O$  emissions from the soils were significantly related to the concentrations of water soluble organic C and N and microbial biomass C and N, respectively. Stepwise regression analysis showed that the emissions of  $CO_2$  and  $N_2O$  from soils and heterotrophic  $N_2O$  emissions could be accounted for by the following equations:

$$y = -(3.6\pm1.2)+(0.023\pm0.003)a+(0.003\pm0.000)b$$
  
+ $(0.68\pm0.22)c$ ,  
 $R^2 = 0.95$ ,  $n = 30$ ,  $P<0.001$ , (1)  
where  $y$  is the soil CO<sub>2</sub> emission (µmol CO<sub>2</sub>-C g<sup>-1</sup> d<sup>-1</sup>)  
and  $a$ ,  $b$  and  $c$  represent the concentration of soil total  
C (mg C g<sup>-1</sup>) and water soluble organic C (µg C g<sup>-1</sup>) and

Table 4 Effects of C<sub>2</sub>H<sub>2</sub> at low concentrations on CO<sub>2</sub> emissions from the soils at the varying distances from trunks

	CO <sub>2</sub> emissions from the	ne soils (µmol CO <sub>2</sub> -C g <sup>-1</sup> dry	I CO <sub>2</sub> -C g <sup>-1</sup> dry soil d <sup>-1</sup> ) after exposure of C <sub>2</sub> H <sub>2</sub> at low concentrations				
Distances from trunks (cm)	0 Pa C₂H₂	10 Pa C <sub>2</sub> H <sub>2</sub>	50 Pa C <sub>2</sub> H <sub>2</sub>	100 Pa C <sub>2</sub> H <sub>2</sub>	LSD <sub>0.05</sub> <sup>a)</sup>		
Broadleaf forest soil							
0	5.05(0.35)	4.92(0.33)	5.13(0.26)	4.89(0.26)	0.91		
50	2.70(0.18)	2.67(0.23)	2.64(0.22)	2.47(0.17)	0.60		
100	2.36(0.26)	2.35(0.24)	2.28(0.20)	2.14(0.25)	0.72		
Coniferous forest soil							
0	3.01(0.99)	2.95(0.95)	2.92(0.95)	2.24(0.69)	2.70		
50	2.01(0.25)	2.00(0.27)	1.88(0.26)	1.85(0.30)	0.80		
100	1.68(0.11)	1.72(0.18)	1.63(0.19)	1.50(0.17)	0.49		
LSD <sub>0.05</sub> b)	1.35	1.31	1.28	1.03			

a) LSD (5%) for comparing differences in CO<sub>2</sub> emissions from each soil after exposure of C<sub>2</sub>H<sub>2</sub> at low concentrations; b) LSD (5%) for comparing differences in CO<sub>2</sub> emissions from different soils after exposure of C<sub>2</sub>H<sub>2</sub> at the same concentration.

**Table 5** Coefficients of correlation between emissions of CO<sub>2</sub> and N<sub>2</sub>O from the soils without C<sub>2</sub>H<sub>2</sub> addition and initial soil properties at the varying distances from trunks<sup>a)</sup>

	N <sub>2</sub> O	CO <sub>2</sub>	Total C	Total N	C:N ratio	DOC	DON	Microbial C	Microbial N	NH <sub>4</sub> -N	NO <sub>3</sub> -N	рН	Hetero N₂O
N <sub>2</sub> O	1												
CO <sub>2</sub>	0.51**	1											
Total C	0.67**	0.92**	1										
Total N	0.78**	0.77**	0.93**	1									
C:N ratio	-0.46**	0.22	-0.01	$-0.33^{*}$	1								
DOC	0.21	0.85**	0.67**	0.47**	0.45**	1							
DON	0.70**	0.83**	0.84**	0.79**	-0.09	0.62**	1						
Microbial C	0.73**	0.67**	0.84**	0.82**	-0.22	$0.40^{*}$	0.68**	1					
Microbial N	0.73**	0.72**	0.85**	0.82**	-0.20	0.44**	0.79**	0.97**	1				
$NH_4^+$ -N	0.25	0.73**	0.75**	0.61**	0.24	0.59**	0.59**	0.61**	0.62**	1			
NO₃N	0.34*	0.53**	$0.38^{*}$	0.30	0.11	0.56**	0.50**	0.21	0.29	0.20	1		
pН	0.32*	0.60**	0.56**	0.54**	-0.10	$0.32^{*}$	0.72**	0.37*	0.50**	0.45**	0.42**	1	
HeteroN₂O	0.37*	0.41*	0.44**	0.36*	0.06	$0.32^{*}$	0.42*	0.49**	0.52**	0.40*	0.12	0.15	1

a) Hetero N<sub>2</sub>O, Heterotrophic N<sub>2</sub>O emission; \* P<0.05, \*\* P<0.01.

soil pH, respectively.

$$y = (24.1\pm10.8)+(8.61\pm2.17)a-(1.72\pm0.50)b$$
  
+ $(0.12\pm0.05)c$ ,

$$R^2 = 0.72, n = 30, P < 0.001,$$
 (2)

where y is the soil N<sub>2</sub>O emission (pmol N<sub>2</sub>O-N g<sup>-1</sup> d<sup>-1</sup>) and a, b and c represent the concentration of soil total N (mg N g<sup>-1</sup>) and exchangeable NH<sub>4</sub><sup>+</sup>-N ( $\mu$ g N g<sup>-1</sup>) and microbial biomass N ( $\mu$ g N g<sup>-1</sup>), respectively, in the soil.

$$y = (9.7 \pm 4.6) + (0.036 \pm 0.011)a,$$
  
 $R^2 = 0.25, n = 30, P < 0.01,$  (3)

where y is the soil heterotrophic  $N_2O$  emission (pmol  $N_2O$ -N  $g^{-1}$   $d^{-1}$ ) and a is the microbial biomass N concentration (µg N  $g^{-1}$ ), respectively, in the soil.

According to the regression models, 95% of the variability in soil CO<sub>2</sub> emissions could be accounted for by the concentrations of soil total C and water soluble organic C and soil pH, and 72% of the variability in soil total N<sub>2</sub>O emissions by the concentrations of soil total N, exchangeable NH<sub>4</sub><sup>+</sup>-N and microbial biomass N and 25% of the variability in heterotrophic nitrification by the soil

microbial biomass N concentration.

There was the nice goodness-of-fit for the regression models of N<sub>2</sub>O and CO<sub>2</sub> emissions from the C<sub>2</sub>H<sub>2</sub>-treated soils (y) against nitrification rate (x), which was described as y = 118.1x-11.9 ( $R^2 = 0.46$ , n = 12, P = 0.01) and y = 3.64x+1.45 ( $R^2 = 0.45$ , n = 12, P = 0.01), respectively. The regression models indicated that both N<sub>2</sub>O and CO<sub>2</sub> emissions from the soils after exposure of C<sub>2</sub>H<sub>2</sub> at low concentrations were positively related to the net nitrification of the soils.

#### 3 Conclusions

The use of  $C_2H_2$  at partial pressures within a range from 10 to 100 Pa  $C_2H_2$  in headspace gas gave a significant decrease in  $N_2O$  emissions from forest soils, and the decrease was the same in each soil after exposure of  $C_2H_2$ . Heterotrophic nitrification could account for 21%-48% of the soil total  $N_2O$  emission; the contribution increased with increasing distances from the coniferous trunks rather than from the hardwood trunks. The

emissions of N<sub>2</sub>O and CO<sub>2</sub> from soils decreased as the distances from trunks increased.

The use of C<sub>2</sub>H<sub>2</sub> at low pressures from 10 to 100 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas did not influence soil microbial biomass N, net N mineralization and microbial respiration, but 100 Pa C<sub>2</sub>H<sub>2</sub> in the headspace could decrease CO<sub>2</sub> emissions from soils. According to the rapid consumption of 10 Pa C<sub>2</sub>H<sub>2</sub> by forest soils and the convenience for laboratory incubations, 50 Pa C<sub>2</sub>H<sub>2</sub> in headspace gas can be used to study the origin of N<sub>2</sub>O emissions from acid forest soils under aerobic conditions and the key associated driving mechanisms.

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N<sub>2</sub>O and CO<sub>2</sub> emissions from forest soils at the varying distances from trunks were associated mainly with the concentrations of soil total C and N, water soluble organic C and N as well as microbial C and N, and both emissions from the C<sub>2</sub>H<sub>2</sub>-treated soils were positively related to the net nitrification of the soils. Soil microbial biomass N concentration accounted for 25% of the variability in heterotrophic nitrification.

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