

## 全球变化背景下多尺度干旱过程及预测研究进展



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**摘要** 干旱是一种周期性的气候异常,主要受气候自然变率驱动,具有发展缓慢、持续时间长、影响范围广等特征。然而,气候变化使得干旱不仅有增加趋势,其特征也在发生变化。例如,发展迅速的干旱,也称骤旱,近年来频繁发生。此外,人类活动通过改变陆地水循环,直接或间接地影响干旱过程。在全球变化背景下,干旱机制研究由海-陆-气相互作用影响气象干旱的气候动力学研究扩展到包含干旱传递过程机理认识、考虑区域尺度人类活动影响,面向农业、水利、生态等行业的农业干旱和水文干旱研究,为认识干旱可预报性、发展预测方法带来了新的挑战。本文将针对多尺度干旱过程及预测,讨论相关的研究进展及未来发展方向。

**关键词** 干旱传递;机制;预测;多尺度;全球变化

干旱主要受气候自然变率驱动,是一种周期性的气候异常。例如,受2015/2016年超级厄尔尼诺事件的影响,西太平洋副热带高压偏东以及异常的欧亚遥相关环流模态导致2015年7—8月东亚夏季风偏弱,使得向北输送的水汽变少,造成我国北方夏季干旱(Wang et al., 2017)。除大尺度海-气相互作用外,陆-气相互作用对干旱持续和加剧也十分重要。2017年3—7月发生在我国东北的春夏连旱,主要由北极涛动异常引起,但陆-气耦合过程维持了贝加尔湖南侧的高压异常,通过准定常的罗斯贝波列影响到下游的东北地区,造成了该地区的持续性干旱(Zeng et al., 2019)。对于更长时间尺度的干旱,如年际和年代际干旱,则主要由大尺度海温异常等引起。比如,赤道东太平洋海温冷异常以及赤道西太平洋和印度洋海温暖异常主导了北半球中纬度地区1998—2002年的干旱(Hoerling, 2003),而东非地区2000—2010年代际干旱与太平洋暖池地区的海温异常密切相关(Lyon and De Witt, 2012)。

相比其他极端事件,干旱具有发展缓慢、持续时间长、影响范围广等特征。在气候变化的背景下,全球和区域尺度的干旱不仅有增加趋势,其特征也在发生变化(黄荣辉等, 1999;符淙斌等, 2005;翟盘茂和邹旭恺, 2005;马柱国和符淙斌, 2006;Sheffield and Wood, 2008; Dai, 2013; Trenberth et al., 2014;张强等, 2015;Huang et al., 2016)。例如,发生在生长季的降水亏缺往往伴随着高温热浪,强烈的太阳辐射导致蒸散发增加、土壤湿度快速降低,从而触发一类发展迅速但强度高、影响大的干旱,也称“骤旱(flash drought)”(Mo and Lettenmaier, 2015;

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2019-11-05 收稿, 2019-11-30 接受

国家自然科学基金资助项目(41875105);国家重点研发计划课题(2018YFA0606002);南京信息工程大学人才启动经费

**引用格式:**袁星,马凤,李华,等,2020.全球变化背景下多尺度干旱过程及预测研究进展[J].大气科学学报,43(1):225-237.

Yuan X, Ma F, Li H, et al., 2020. A review on multi-scale drought processes and prediction under global change[J]. *Trans Atmos Sci*, 43(1): 225-237. doi: 10.13878/j.cnki.dqkxxb.20191105005. (in Chinese).

Yuan et al., 2015b; Wang et al., 2016)。其中,美国中部大平原 2012 年骤旱造成了 120 亿美元经济损失(Hoerling et al., 2014),我国南方 2013 年夏季骤旱影响了 13 个省,仅湖南、贵州两省受灾作物面积就达 200 万公顷(<http://www.chinaam.com.cn>)。未来骤旱的风险很可能由于持续升高的全球气温而进一步增加(Yuan et al., 2019)。

干旱的发生受不同时空尺度因子的影响,其复杂的发生和恢复机制,使其成为最复杂的极端事件之一(Yuan and Wood, 2013; Kiem et al., 2016)。根据美国气象学会的定义,干旱通常分为气象干旱、农业干旱、水文干旱和社会经济干旱(Wilhite, 2000; Mishra and Singh, 2010)。气象干旱通常起源于气候异常(图 1),如降水不足,高温异常等(Cook et al., 2014; Livneh and Hoerling, 2016; Luo et al., 2017)。气候异常可能会进一步造成蒸发增加,土壤湿度降低,径流量减少,导致农业干旱和水文干旱,从而影响社会经济的发展。在全球变化背景下,土地利用变化、灌溉和水库调度等人类活动,通过改变水文循环过程,影响干旱的发生和发展,使干旱已不再是单纯的气候异常。

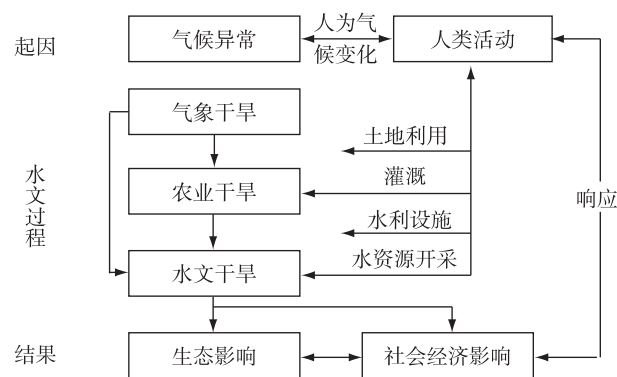


图 1 干旱类型及传递过程

Fig.1 Drought types and the propagation processes

由于干旱成因的复杂性及多尺度时空特征,干旱预测面临着巨大的挑战。干旱预测通常基于 3 种方法,即统计预测、动力预测和统计动力混合预测方法(Mishra and Singh, 2011; Mariotti et al., 2013; Pozzi et al., 2013)。统计预测方法基于历史资料,通过相关性分析,将与干旱联系密切的影响因子作为预测因子进行预测。此外,随着人们对气候系统认识的提高和计算能力的进步,基于先进的气候系统模式进行动力预测已成为目前干旱预测的主流方向(Yuan et al., 2015c)。然而,统计预测和动力预测都有其各自的优缺点,因此出现了统计和动力混

合预测方法(Schepen et al., 2014; Schepen and Wang, 2015; Bennett et al., 2016),在近十几年得到了较快发展。本文主要对次季节到季节尺度的多尺度干旱过程及预测方面的研究进展进行回顾,讨论目前研究面临的挑战和未来发展方向。

## 1 多尺度干旱过程及机制

### 1.1 干旱分类和干旱指数

基于物理和社会经济因素,通常将干旱分成气象、农业、水文和社会经济干旱四种类型(Mishra and Singh, 2010)。当然,还有研究进一步细化为生态干旱、地表水干旱、地下水干旱等。气象干旱主要由长时间降水不足引起。常用的气象干旱指数包括标准化降水指数(Standardized Precipitation Index, SPI; McKee et al., 1993)、标准化降水蒸发指数(Standardized Precipitation Evaporation Index, SPEI; Vicente-Serrano et al., 2010)、帕尔默旱度指数(Palmer Drought Severity Index, PDSI; Palmer, 1965; Dai, 2011)和修正的 PDSI 指数(Liu et al., 2017)等。其中, SPI 由于计算简单,可适用于多时间尺度,得到广泛应用(袁文平和周广胜, 2004; Hayes et al., 2011; Yuan and Wood, 2013; Ma et al., 2015; Schneider et al., 2015)。农业干旱通常与土壤水分不足有关,影响植被和作物生长。最为常用的农业干旱指数是土壤湿度分位数(Soil Moisture Quantile, SMQ; Sheffield, 2004),当 SMQ 连续低于某一分位数阈值(如 20%)时,则会发生农业干旱。由于土壤湿度观测资料的缺乏,通常采用陆面水文模型模拟(Sheffield, 2004; Wang et al., 2011; Wu et al., 2011; Xia et al., 2014)或卫星遥感反演(Liu et al., 2012; Pan et al., 2012; Yuan et al., 2015b; Bi et al., 2016)等方法获得土壤湿度数据。水文干旱表现为地表径流或水库水资源减少,地下水位下降等。水文干旱与其对生态环境的影响直接相关,近年来得到更多关注。常用的水文干旱指数包括标准化径流指数(Standardized Runoff Index, SRI; Shukla and Wood, 2008; 任立良等, 2016)、标准化流量指数(SFI; Yuan et al., 2017)和水库水位(Hayes et al., 2011)等。社会经济干旱涉及社会经济利益,如人体健康、农业产量、水供给量、森林生产量和森林火灾等。社会经济干旱没有统一的指数,通常用作物产量、水短缺量和水质来表征(Quiring and Papakryiakou, 2003; Mosley, 2015)。虽然干旱指数众多(Zargar et al., 2011; 张强等, 2011; 沈彦军等,

2013),但单一的干旱指数并不能反映多种类型的复杂干旱特征。近 10 a 来,通过结合一系列水文气象变量或干旱指数,多干旱指数或综合干旱指数得到发展和应用(Hao and Singh, 2015; Qin et al., 2015)。如美国干旱监测系统(USDM, Svoboda et al., 2002)利用多种气象水文指标、模型模拟和遥感并基于当前资料与历史条件的比较来分析干旱状况,是综合干旱指标较为成功的案例。此外,还有国家气候中心发展的综合气象干旱指标 CI(张强等, 2006)等。

## 1.2 季节干旱机制及传递过程

季节性气象干旱通常是由海温或其他遥相关型异常引起大尺度环流异常(如高压异常)导致(Dai, 2011; Lu et al., 2011; 张强等, 2011)。我国学者在气象干旱成因方面做了大量诊断分析工作,逐步认识到海温异常强迫、大气环流内部异常如副热带高压和中纬度阻塞高压的强度和位置异常、气候年代际变化等因素对于干旱发生的影响(毕慕莹和丁一汇, 1992; 李维京等, 2003; 杨修群等, 2005; 陶诗言等, 2009; 琚建华等, 2011; 封国林等, 2012; 钱维宏和张宗婕, 2012; Sun and Yang, 2012; Jin et al., 2013)。

厄尔尼诺-南方涛动(ENSO)可以影响全球许多地方的季节性气候,包括美国南部和北部、非洲东部和南部、亚洲和澳洲等地区(Smith et al., 2012; Schubert et al., 2016)。研究发现,ENSO 通过影响西太平洋副高和季风来影响我国气候(李栋梁和姚辉, 1991; 张人禾, 1999; Zhang et al., 1999; Wang et al., 2000; Zhai et al., 2016)。通常,当厄尔尼诺现象发生时,赤道中东太平洋海温出现正异常,热带沃克环流出现异常,并通过海气相互作用等途径引起西北太平洋反气旋,西太平洋副高加强,导致传输到我国的水汽出现异常,易使我国北方地区出现干旱而长江流域发生洪涝。同时,印度洋海温的电容器效应会延长 ENSO 对我国夏季气候的影响。研究表明 20 世纪 80 年代末以来,冬季 ENSO 的信号会储存在印度洋中并在夏季导致印太地区的海温异常,进而直接影响我国西北地区的干旱异常(Zhu et al., 2019)。2015 年夏季我国华北地区出现的严重干旱现象就与 2015/2016 强厄尔尼诺现象有密切的联系(Zhai et al., 2016; Wang et al., 2017; Ma et al., 2018)。然而,强厄尔尼诺并不意味着我国北方夏季极端干旱发生(Wang et al., 2017)。除 ENSO 之外,其他遥相关信号(如印度洋海温异常、太平洋年代际振荡 PDO 等)也对我国气候异常有显著影响

(吴统文和钱正安, 1996; 马柱国, 2007)。2014 年夏季,我国出现了典型的“南涝北旱”事件,研究表明北太平洋海温异常激发的日本-太平洋遥相关型和印度洋海温异常激发的丝绸之路遥相关型对此次事件发挥了决定性的作用(Wang and He, 2015; Xu et al., 2017a; Xu et al., 2018)。2014 年的干旱事件还存在季节内北移的现象,主要由东亚急流的北移造成(Xu et al., 2017b)。随着全球变暖的加剧,北极地区经历了更为显著的增暖,导致北极海冰的快速减少,从而对我国的气候异常产生显著影响。研究表明,巴伦支海春季海冰的减少会加剧我国东北地区的夏季干旱(Li et al., 2018)。一些大气遥相关型,例如日本-太平洋遥相关型、丝绸之路遥相关型、欧亚遥相关型和北极-欧亚遥相关型会直接影响局地的高压和下沉等导致干旱事件的发生(Wang and He, 2015; Li et al., 2018)。

此外,局地气候因素(如地形)、陆气耦合(如土壤湿度降低,蒸散发降低,气温升高)和人为气候变化也会影响干旱的发展、持续和强度(辛晓歌等, 2009; Dai, 2011; Wang et al., 2015; Van Loon et al., 2016)。众多研究表明陆面和大气之间水分和能量的反馈对于干旱的强度和持续时间影响显著(Hong and Kalnay, 2000; Schubert et al., 2004; 陈海山和周晶, 2013; Yang et al., 2016)。李崇银等(2013)利用区域模式研究了前期土壤湿度对 2009/2010 年冬季云南干旱的影响,发现前期土壤湿度减少一半可使冬季降水量平均减少 30% 以上,部分地区甚至减少 50%; Liu et al. (2014)研究了东亚地区土壤湿度-降水之间的反馈,发现土壤湿度异常对中、高纬度地区降水影响显著。干旱过程中的极端高温通过陆气相互作用将进一步加剧干旱,但陆气相互作用相对于海气相互作用在干旱过程中具体贡献如何仍需进一步研究。

由气象干旱演变到农业干旱和水文干旱的过程称为干旱传递过程。目前已有不少干旱传递方面的研究(Van Loon et al., 2012; Huang et al., 2017; Yuan et al., 2017; Wu et al., 2018a; Ma et al., 2019b)。一般认为,除地下水丰富地区(Peters et al., 2003, 2006)和积雪地区外,水文干旱与气象干旱之间存在明显的联系(Zhao et al., 2016)。一般来说,存在 4 种干旱传递过程(图 2, Van loon, 2013): 1) 气象干旱合并为水文干旱(pooling); 2) 在气象干旱开始阶段,如果流域水储量较高,水文干旱会受到抑制(attenuation); 3) 水文干旱滞后于气象干旱(lag); 4) 由

气象干旱到农业干旱和水文干旱,干旱持续时间逐渐增加(*lengthening*)。气候特征对干旱传递过程有直接的控制作用,其中降水的时机对干旱传递的影响最大(Apurv et al.,2017)。在气象干旱传递到农业或水文干旱的过程中不仅受到气候特征的影响,还与陆地水文特征有关,如流域储水量(土壤水、地下水、积雪等)、地质特征(岩石类型和硬度)、地形、土壤特征(土壤质地和结构等)、排水系统、土地利用和植被等。然而,各个因子的影响机制和程度尚不明确。Apurv et al.(2017)基于地下水的补给特征,将干旱传递机制概括为 3 种类型:1)干旱期地下水存在季节性的持续补给,导致水文干旱持续时间比气象干旱短;2)干旱期地下水补给受到抑制,导致水文干旱持续时间长于气象干旱;3)地下水补给缺乏季节性变化,主要受降水控制,导致水文干旱持续时间与气象干旱接近。

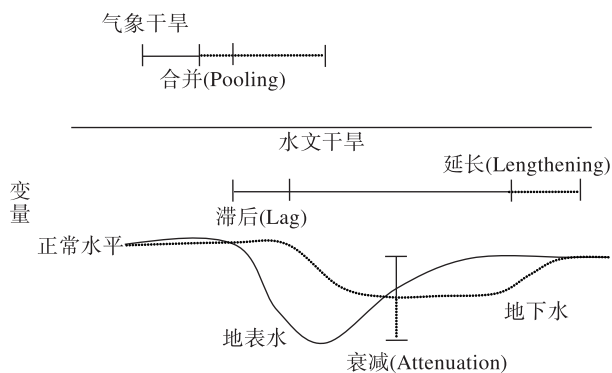


图 2 气象干旱到水文干旱传递特征示意

Fig.2 Diagram for the propagation of meteorological drought to hydrological drought

此外,人类活动如土地利用变化、灌溉活动和水库调度等也会通过水文循环对干旱传递过程产生影响。其中,土地利用变化通过影响产汇流机制影响干旱的传递过程(Zhou et al.,2019)。水库调度和灌溉会改变水文干旱对气象干旱的响应关系,且对于上下游的水文干旱具有不同的影响(Wu et al.,2016;Yuan et al.,2017;Wu et al.,2018b;Ma et al.,2019a)。此外,研究者还基于模型试验量化了不同地区气候变化和人类活动对水文干旱及干旱传递的影响和贡献(Zhang et al.,2018)。研究发现,两者的相对贡献随区域和季节变化,在人类活动影响较大的地区,如黑河中游,人类活动对灌溉季水文干旱的贡献率可达 89% 以上(Ma et al.,2019a)。

### 1.3 骤旱

在全球变暖的背景下,干旱和热浪同时发生的概率明显增加(Hao et al.,2013;Chiang et al.,

2018),从而触发一类发展迅速但强度高、影响大的干旱,也称骤旱。该事件一般发生在次季节尺度(几周到几个月不等)。与传统的干旱类似,通常伴随高压异常和大气下沉运动(Ford and Labosier,2017),多发生在季节干旱开始或结束的过渡时期(Wang and Yuan,2018a)。近年来,骤旱频繁发生,如 2012 年美国中部骤旱(Hoerling et al.,2014)、2013 年中国长江中下游骤旱(Yuan et al.,2015b)、2015/2016 年南部非洲骤旱(Yuan et al.,2018)和 2017 年美国北部骤旱(Gerken et al.,2018)等。尽管相关研究逐渐增多,然而对于骤旱仍没有统一的定义。一部分研究基于干旱演变过程中气温、蒸散发和土壤湿度的变化,将骤旱分为两种类型,一种是由高温引发,表现为高温异常使蒸散发迅速增加,土壤湿度降低,主要发生在湿润地区如我国南方地区;另一种是由降水缺乏引发的,表现为降水不足导致蒸散发降低,气温升高,主要发生在半干旱地区如我国北方地区(Mo and Lettenmaier,2015,2016;Wang and Yuan,2018a)。两种定义主要基于热浪和土壤水分不足共同发生的情况,但并未刻画骤旱的快速发展过程。最近,Yuan et al.(2019)考虑土壤湿度的快速降低和干旱的持续性,将骤旱定义为 5 d 平均的土壤湿度从 40% 分位数以上迅速下降到 20% 分位数以下,下降速率不低于 5%,并且至少持续 15 d(图 3)。结果发现,新的定义可以同时反映干旱程度及其发展迅速的特点(Yuan et al.,2019),而以前的定义则高估了骤旱发生频率,使骤旱的持续时间过短,从而可能不会对生态系统产生显著影响。

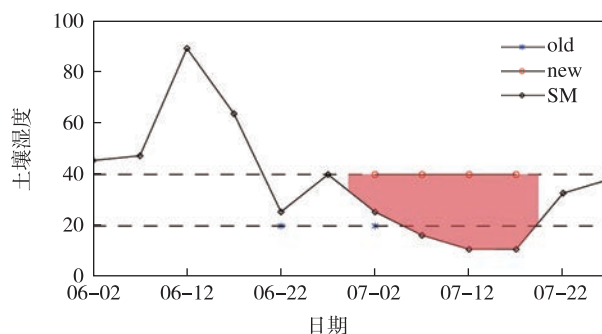


图 3 基于土壤水分下降速率和干旱持续时间的骤旱事件识别方法(引自 Yuan et al.,2019)

Fig.3 Definition of a flash drought event based on rapid decline in soil moisture and drought duration (from Yuan et al.,2019)

在全球变暖的背景下,骤旱发生的频率也在增加。1979—2010 年我国的骤旱事件平均增加了 109%,主要归因于气温的长期变暖趋势(Wang et

al.,2016)。因此,一些研究开始关注人为气候变暖对骤旱发生的影响和贡献。研究发现,人为气候变化使非洲南部的骤旱风险在过去 60 a 增加了 3 倍 (Yuan et al.,2018),温室气体排放增加导致的人为气候变暖能解释中国地区 77% 的骤旱增加趋势 (Yuan et al.,2019)。此外,人口增长也使骤旱的风险增加。在未来气候变化情景下,尽管我国生长季降水有所增加,但由于人为排放温室气体增加导致气候变暖,蒸散发增加,土壤湿度减少,许多地区骤旱的频率仍会继续增加,如我国南部湿润地区 (Yuan et al.,2019)。

## 2 干旱预测和可预报性

### 2.1 季节干旱可预报性

认识干旱可预报性是提高干旱预报技巧的重要前提。气象干旱可预报性的来源主要包括外强迫异常,如海温异常等(图 4)。海温异常的准确观测和预测明显提高了气象干旱的预报能力,尤其是在遥相关较强的地区 (Hoerling, 2003; Schubert, 2004; 王蕾和张人禾, 2006; Kallis, 2008; Yuan and Wood, 2013; Seager and Hoerling, 2014; Schubert et al., 2016)。例如,赤道中东太平洋海温异常导致的 ENSO 现象是全球许多地区干旱最重要的可预报性来源。此外,局地气候因素如青藏高原大地形等也是区域季节气候可预报性的来源之一 (吴统文和钱正安, 1996; Liu et al., 2007; Gao, 2014)。气象干旱可预报性的定量评估主要基于模型最优假设,即一个预报系统能够达到的最大预报能力 (Luo and Wood, 2006; Ma et al., 2015, 2018; Wang and Yuan, 2018b)。通常将模式的一个集合成员作为观测,其他集合成员的平均作为预测进行相关性计算得到。研究发现,气象干旱的可预报性通常要高于实际预报能力,说明预报存在一定的可提升空间 (Ma et al., 2015)。此外,区域陆表特征 (如土壤湿度、植被和雪盖等) 的异常也会加剧或减弱干旱强度,影响干旱的发展 (Yuan and Wood, 2013; Hao et al., 2018),从而通过陆气耦合机制为提高气候预测提供有效信息 (Van den Hurk et al., 2012)。

相对于气象干旱,陆面水文初始信息是农业和水文干旱的重要可预报性来源(图 4)。其中陆面水文初始信息包括积雪、土壤湿度、地下水等初始异常。水文可预报性的研究主要集中在定量评估气象强迫和水文初始信息对水文干旱的相对贡献 (Yuan et al., 2016)。较为常用的方法主要基于两组模型

模拟试验(图 5):1)集合径流预报试验 ESP,即根据流域当前水文初始状态,采用历史同时期不同年份气象强迫作为驱动水文模型的输入集合,得到流域的水文过程;2)逆集合径流预报实验 revESP (Wood and Lettenmaier, 2008; Shukla and Lettenmaier, 2011),即将历史同时期不同年份水文初始状态作为集合,采用当前实际气象强迫驱动水文模型,得到流域的水文过程。然后利用 ESP 和 revESP 试验的均方根误差比值来评估气象强迫和水文初始信息的相对重要性。

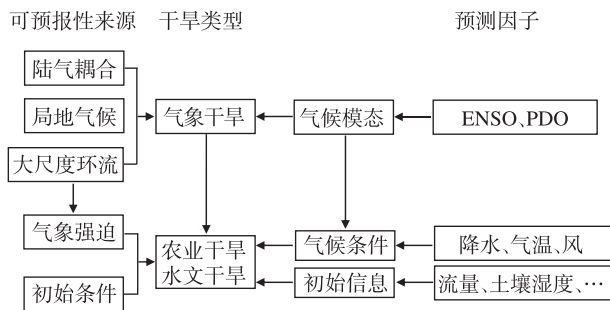


图 4 干旱可预报性来源和预测因子

Fig.4 Sources of drought predictability and predictors

研究发现,两者的相对贡献随区域、季节、预见期而变化。随着预见期的增加,水文初始信息的贡献降低,而气象强迫的贡献逐渐增加,且随季节变化。如在我国黄河流域,水文初始信息对径流预报的贡献在冷季和干季要高于气象强迫的贡献,预见期可达 2~5 mon (Yuan et al., 2016)。又如在我国黑河流域,冷季和暖季水文初始信息的贡献高于气象强迫的预见期在 2~7 mon,而在暖季和湿季,预见期则不超过 1 mon (Ma et al., 2018)。

### 2.2 季节干旱预测

季节干旱预测对提前采取抗旱减灾措施、保障水资源和生态系统可持续发展具有十分重要的意义。干旱预测一般采取 3 种方法:统计预测、动力预测和统计动力混合预测方法 (Hao et al., 2018)。统计预测方法 (如回归模型等) 主要基于历史资料,通过相关性分析,将与干旱联系密切的影响因子作为预测因子,建立预测因子与干旱之间的统计或物理模型进行预测 (Yuan et al., 2015c)。该方法的优点是计算简单,通常用作动力预测的参考标准,并提供辅助预测信息 (Hao et al., 2017)。

动力预测主要依赖于全球天气或气候模式的季节预测产品。这些气候模式是基于大气、海洋、冰冻圈和陆地圈的物理过程刻画并相互耦合发展起来的 (Quan et al., 2012; Yoon et al., 2012; Yuan and

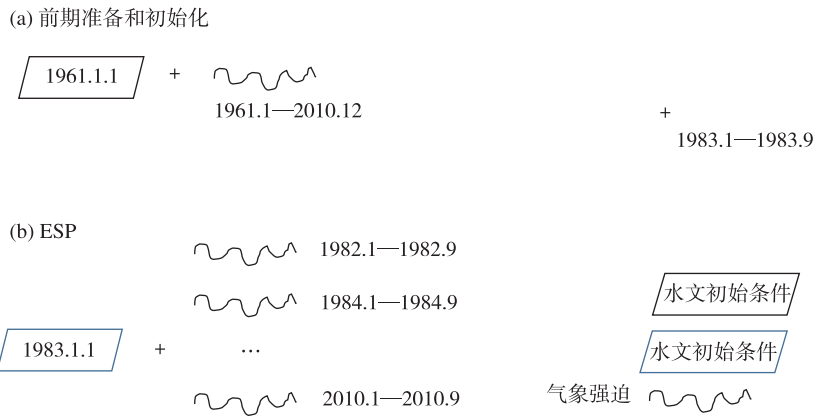


图 5 ESP 和 revESP 试验示意

Fig.5 Illustration of the ESP and revESP experiments

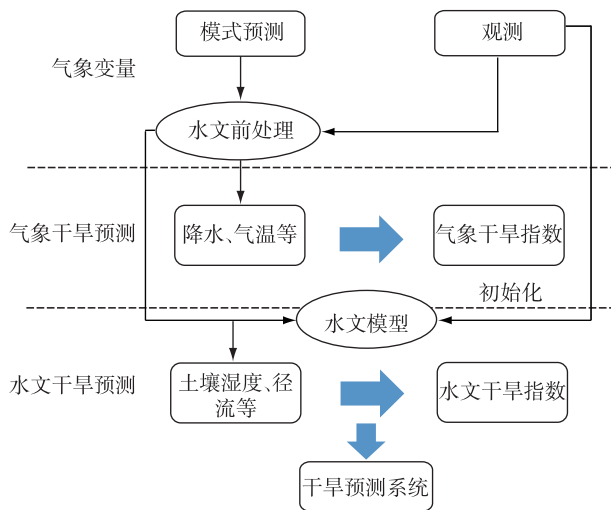


图 6 干旱动力预测系统示意

Fig.6 Illustration of the dynamical drought forecasting system

Wood 2013; Schubert et al., 2016)。目前比较流行的有美国国家环境预报中心 NCEP 的气候预测系统 CFS、欧洲中期天气预报中心 ECMWF 的 IFS 等。它们使得干旱预测可以在区域或全球尺度上进行。在预测农业干旱和水文干旱时,还需要采用陆面水文模型将气候异常信息传递到水文变量中(图 6, Wood et al., 2002; Yuan et al., 2013; Thober et al., 2015)。在气候预测模式得到广泛应用之前,水文干旱预报采用的是集合径流预报方法 ESP,该方法将历史气象观测集合作为未来预报驱动水文模型进行预测。因此,其预报能力十分依赖于水文初始状态,且预报精度偏低,目前多用其作为气候模式驱动水文模型进行预测的参考标准(Yuan et al., 2013, 2015b; Ma et al., 2018)。随着海陆气耦合模式的不断发展,利用耦合模式预报产品驱动水文模型进行

农业和水文干旱预报得到广泛应用(Luo and Wood, 2007; Li et al., 2008; Yuan et al., 2013; Sheffield et al., 2014; Yuan et al., 2015b, 2015c; Ma et al., 2018; Yao and Yuan, 2018)。常用的水文模型包括 VIC(Liang et al., 1996)、分布式时变增益水文模型 DTVGM(夏军等, 2004)、SWAT 和新安江模型等。

动力预测的优点是基于物理过程,精确度较高,但是计算比较复杂,同时也会带来大量的不确定性,如模型输入、模型参数、水文初始信息的不确定性等。由于模型分辨率和系统误差的限制(Roundy and Wood, 2015),首先需要对气候模式的气象要素预报产品进行前处理,如偏差校正和降尺度(Wood et al., 2002)、典型事件分析(Roundy and Wood, 2015)和集合前处理(Yuan et al., 2013; Ye et al., 2017)等,以便减少陆面水文模型的预报误差。利用多气候模式/水文模型集合预测是减少不确定性,提高预报能力的主要手段之一(Yuan et al., 2015b; Xu et al., 2017a; Yao and Yuan, 2018)。此外,通过同化技术提高观测和初始信息的精度也可以提高农业和水文干旱的预报能力。由于统计和动力预测具备各自的优缺点,两种方法结合或混合预测方法在干旱预警和预测系统中得到发展和应用。混合预测方法主要是对气候模式预报进行偏差校正,然后结合统计模型对得到的多源预测结果进行融合。比较常用的融合方法包括回归模型、贝叶斯后验分布和贝叶斯平均 BMA 等。目前基于 CBaM (Calibration, Bridging, and Merging) 的统计-动力混合预测方法已经被用于提高气候变量和干旱的预测(Schepen et al., 2014; Schepen and Wang, 2015; Bennett et al., 2016)。

随着干旱预测方法的不断发展,干旱预警和预测系统在区域和全球尺度上得到建立和应用(Yuan et al., 2015b)。例如,美国干旱监测图于1999年完成并上线发布,基于多干旱指标和决策信息等,每周发布干旱预警信息,为政府决策、农业生产等提供重要指引(Svoboda et al., 2002)。Yuan et al. (2015b)利用NMME多气候模式的集合预报产品驱动VIC水文模型建立了全球季节水文预测系统,用于季节水文和极端事件的预测。该系统与ESP相比,可以更好地监测和预测土壤干旱,并提供更加准确的水文干旱预测信息。此外,欧洲干旱观测系统(<http://edo.jrc.ec.europa.eu/edov2/php/index.php?id=1000>)、英国水文趋势预测系统(<http://www.hydotuk.net>)也得到广泛应用。我国国家气候中心以综合气象干旱指数为标准,分别发布全国逐日和旬干旱综合监测预警(<http://cmdp.ncc-cma.net/en/>)。2015年我国科技部批准了“干旱气象科学研究-我国北方干旱致灾过程及机理”项目(DroughtEX\_China),目的是利用多源观测和多尺度模型为多个部门(尤其是农业部门)提供早期干旱预警和预测信息(Li et al., 2019)。

### 3 挑战与展望

尽管干旱研究不断受到各国关注,但由于干旱的复杂性,干旱过程的深入理解以及干旱事件的可靠预测仍面临很大挑战,需要进一步理解不同尺度和不同类型干旱的发生机制和可预报性。本文对于旱过程和预测研究提出如下几点展望:

1) 水文气象观测数据(站点观测、遥感、再分析等)是干旱问题研究的基础。然而,站点观测存在区域代表性差、覆盖率低等缺陷,遥感监测受其他因素(如云、观测平台和内在误差等)的影响,数据产品质量较低或不连续(Sheffield et al., 2014)。此外,再分析等模型模拟资料也存在一定的限制,例如模型需要大量的参数,并且参数会带来很大的不确定性。因此迫切需要提供更高精度的长期连续或实时的数据产品,为干旱过程研究和干旱预警预测提供数据支撑。如何利用先进的技术将多源数据进行融合,以提高数据的精度和分辨率仍需要进一步探索。

2) 目前干旱指数较多,但缺乏一个国际普遍认可的干旱指数,尤其是农业和水文干旱指数。大多数干旱指数仅描述单一类型的干旱,在分析干旱过程和建立干旱预测系统时,需要采用多干旱指数或

综合干旱指数。针对特定的应用部门(农业、林业、用水等部门),如何将众多的信息结合干旱影响发展成与其直接相关的指数仍面临巨大挑战。此外,针对气候变化背景下出现的新现象(如骤旱),也亟需利用合适的指数,开展相应的监测和预警研究,以提高气候变化的适应能力。

3) 农业和水文干旱不仅受气候异常影响,很大程度上也受到陆面水文过程和人类活动的影响。Yuan et al. (2017)发现,由于黄河流域径流受到人为用水活动的显著影响,该地区的水文干旱预测很大程度上依赖于对人为用水活动的描述。换言之,人类用水活动对黄河流域水文干旱可预报性的影响不亚于气候变率(Yuan et al., 2017)。但目前大部分陆面水文模型对人类活动描述的不确定性较大,如何在高分辨率架构下将不同类型人类活动协调地加入陆面水文模型中并用于干旱预测,将是未来很长一段时间研究的重点。此外,物理模型结合人工智能等统计学习方法,将是未来预测研究很有生命力的一个方向,特别是如何将干旱预测结果更加智能地转化为有效的干旱预警并指导应对措施制定。

4) 尽管气候预测取得了长足的进步并被广泛用于农业和水文干旱预测,陆面水文过程的初始记忆性始终是水文预测的重要可预报性来源。过去的研究集中在探讨地表水、表层和根区土壤水异常对水文干旱预报的影响,但对更深层土壤水甚至地下水记忆性的认识仍然十分薄弱。随着GRACE重力卫星反演陆地水储量数据、SMAP等微波遥感反演土壤湿度数据的广泛应用,结合即将发射SWOT地表水和海洋地形卫星提供的高精度地表水数据,有望进一步推动陆地记忆性研究。如何分离地表水(包括积雪)、土壤水、地下水等对农业和水文干旱可预报性的相对贡献,将是未来的一个研究热点和难点。

### 4 结论与讨论

本文对次季节到季节尺度干旱过程及预测方面的研究进展进行了简要回顾。气候异常引发气象干旱,发展到一定程度会导致农业和水文干旱,对社会、经济和环境产生影响。大尺度气候模态是气象干旱可预报性的重要来源。除此之外,陆气耦合和人为气候变化也会影响干旱的发展。在全球变暖背景下,次季节尺度的干旱(骤旱)频率也在不断增加,这种干旱发展迅速,通常会伴随热浪的发生。气

象到农业或水文干旱的传递过程不仅受气候条件的影响,还很大程度受陆气相互作用、陆地人类活动的影响。

尽管过去在干旱机制研究等方面取得了重大进展,干旱预测仍面临巨大挑战。季节干旱的预报能力主要限制在与 ENSO 相关性较强的热带和亚热带地区,或陆面特征记忆力较强的中纬度地区 (Yuan et al., 2015c; Hao et al., 2018)。由于气候系统的混沌特性,长预见期的跨季节预测能力较低,从而也会影响干旱的预报能力。此外,自然和人为的

气候变化改变水文循环过程,为干旱的预测带来新的挑战。因此,进一步加深对不同干旱发生和发展机制的认识、通过数据同化技术提高观测数据精度、精细参数化方案提高模型模拟自然和人为活动影响的精度、发展综合干旱指数和通过概率预报降低不确定性是提高干旱预报能力的关键。最后,干旱问题研究要与干旱影响相结合,如森林火灾、植被生态和粮食作物产量等,才能更好地为各部门服务。

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## A review on multi-scale drought processes and prediction under global change

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Drought is defined as a period of climate anomaly, which is mainly driven by natural climate variability. It has the characteristics of slow development, long duration and wide range. However, due to climate change, drought is occurring with increased frequency and changing characteristics. For example, a type of drought with rapid development, also known as flash drought, has occurred frequently in recent years. Moreover, human activities also affect the processes of drought by altering the terrestrial water cycle. Under global change, drought research has been extended from understanding climate dynamics of meteorological drought by considering ocean-land-atmosphere interaction to investigating the propagation from meteorological drought to agricultural and hydrological droughts with the influence of human activities. The latter will provide direct climate service for agricultural, water resources, forest sectors, but they also bring new challenges for understanding of drought predictability and developing drought forecasting methods, especially in the Anthropocene. In this paper, we review the advances in multi-scale drought processes and prediction, and discuss future directions.

**drought propagation; mechanism; prediction; multi-scale; global change**

doi:10.13878/j.cnki.dqkxxb.20191105005

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